Prediction and Migration of Surface-related Resonant Multiples

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SUMMARY

Surface-related resonant multiples can be migrated to achieve better resolution than migrating primary reflections. We now derive the formula for migrating surface-related resonant multiples, and show its super-resolution characteristics. Moreover, a method is proposed to predict surface-related resonant multiples with zero-offset primary reflections. The prediction can be used to identify and extract the true resonant multiple from other events. Both synthetic and field data are used to validate this prediction.

INTRODUCTION

A first-order surface-related resonant multiple is a zero-offset multiple which enjoys two round trips between the free-surface and the common reflector. As shown in Figure 1, the raypaths are coincident with one another. Resonant multiples can be migrated to achieve better resolution than migrating primary reflections. In fact, the imaging resolution can be better than half of the wavelength so that this is considered super-resolution imaging (Schuster and Huang, 2014).

This report firstly derives the migration operator for surface-related resonant multiples, and explains why this operator shows super-resolution characteristics. Then, we present the theory of using zero-offset (ZO) primary reflections to predict resonant multiples. This data-based prediction is used to identify and to extract the true resonant multiples in the data. This extraction is required before migration, because the resonant multiples are entangled with other multiple and primary events. Migrating non-resonant multiple events with the resonant multiple migration operator will generate artifacts in the migration image.

After discussing the theory section, synthetic and field data results are shown, and the conclusions are drawn in the last section.

THEORY: MIGRATION OPERATOR OF RESONANT MULTIPLES

The Lippmann-Schwinger equations says that the forward modeling formula for a 1st-order resonant multiple from the free surface (Raypath is shown in Figure 1.) can be approximated by

$$d(\mathbf{g}) = \int_V \omega^4 \int_B [G(\mathbf{g}|\mathbf{x})^D]^2 G(\mathbf{x}|\mathbf{x})^M m(\mathbf{x}) dx^3 d\omega, \quad (1)$$

where V is the set of points in model space and for convenience we assume a wide band source so the wavelet spectrum $W(\omega) = 1$. The term $G(\mathbf{x}|\mathbf{g})^D$ is the direct wave Green's

function for a source at $\mathbf{g}$ and receiver at $\mathbf{x}$ in a heterogeneous medium that honors the 1-way path $\overline{\mathbf{x}\mathbf{g}}$ in Figure 1; and $G(\mathbf{x}|\mathbf{x})^M$ is the 1st-order multiple Green's function that honors the two-way $\overline{\mathbf{x}\mathbf{g}'\mathbf{x}}$ path in Figure 1 for a source at $\mathbf{x}$ and receiver at $\mathbf{x}$. The recorded 1st-order resonant data are represented by $d(\mathbf{g})$ for a source and receiver at $\mathbf{g}$ for upcoming pressure waves just below the free surface. Here, $m(\mathbf{x})$ represents the squared reflectivity at the location $\mathbf{x}$ modulated by a -1 amplitude associated with the reflection from the free surface. The reflectivity at $\mathbf{x}$ is squared because the resonant multiple bounces twice from the reflector at $\mathbf{x}$.

Figure 1: Raypath of a 1st-order resonant multiple in a homogeneous medium with a dipping interface. Star and triangle represent the source and receiver, respectively. The red arrows represent the ray which starts from the source location $\mathbf{g}$, hits $\mathbf{x}$ at the dipping interface, bounces vertically up and hits $\mathbf{g}'$ at the free surface boundary. Then the ray repeats its path by going vertically down, reflecting again at $\mathbf{x}$ and is finally recorded by the receiver at $\mathbf{g}$.

The migration formula is the adjoint of equation 1 applied to the data:

$$m(\mathbf{x})^{mig} = \int \omega^4 \int_B \{ [G(\mathbf{g}|\mathbf{x})^D]^2 G(\mathbf{x}|\mathbf{x})^M \}^* d(\mathbf{g}) dx_g^2 d\omega, \quad (2)$$

where the integration is over the 2D recording survey (with locations denoted by the set of points in B) just below the free surface. Here, $m(\mathbf{x})^{mig}$ is the migration image of the resonant multiples.

Under high frequency asymptotic assumption (Bleistein et al., 2001), $G(\mathbf{x}|\mathbf{x})^M = A(\mathbf{x}, \mathbf{x})^M e^{i2\omega\tau_{xx}^{min}}$, $G(\mathbf{x}|\mathbf{g})^D = A(\mathbf{g}, \mathbf{x}) e^{i\omega\tau_{xg}}$, and $d(\mathbf{g}) = A(\mathbf{g}) e^{i\omega\tau_g}$. τ_{xx}^{min} is the shortest traveltimes of the ray $\overline{\mathbf{x}\mathbf{g}'\mathbf{x}}$, which in practice can be calculated by a plane-wave eikonal solver. τ_{xg} is the traveltime of the direct arrival from $\mathbf{g}$ to $\mathbf{x}$. τ_g is the traveltime of the 1st-order resonant multiples. Neglecting the amplitude parts, the kernel in equation 2 can be approximated by

$$\{ [G(\mathbf{g}|\mathbf{x})^D]^2 G(\mathbf{x}|\mathbf{x})^M \}^* d(\mathbf{g}) \approx e^{i\omega(\tau_g - 2(\tau_{xx}^{min} + \tau_{xg}))}. \quad (3)$$

For a fixed τ_g , $\tau_g - 2(\tau_{xx}^{min} + \tau_{xg}) = 0$ defines an isochron in the model space, which in a homogeneous medium can be analytically formulated as

$$z = \frac{t_m v}{4} - \frac{(x - x_g)^2}{t_m v}, \quad (4)$$

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where v is the velocity of the homogeneous medium, and z and x are the vertical and horizontal component of $\mathbf{x}$, respectively. This isochron is shown as the red curve in Figure 2a. If the traveltimes of a resonant multiple varies from $t_g - T_0/2$ to $t_g + T_0/2$, where T_0 is the dominant period of the resonant multiple wavelet, the isochron changes correspondingly from the inner to the outer dashed blue line in Figure 2a. Between the two lines is the migration response of a single resonant multiple wavelet, which is defined as the migration operator of resonant multiples. Similar analysis can be applied to the ZO primary reflection case, and its migration operator is shown in Figure 2b. The comparison between Figures 2a and 2b shows the resonant multiple operator is skinnier and so has a better resolution than the ZO primary reflection operator. For example, for the image point directly below the source/receiver position ($x = x_g$), the thickness of the resonant multiple operator is a quarter of the wavelength, which is half compared to the primary reflection operator. In other words, resonant multiple migration shows subwavelength (also called super-resolution) characteristics.

Figure 2: Single-trace migration operators for (a) resonant multiples, and (b) primary reflections.

A synthetic experiment is designed to illustrate super-resolution imaging achieved by migrating resonant multiples. As shown in Figure 3, 16 point scatterers are placed in the model. In Figure 3a, diffractions of the point scatters are generated by Kirchhoff modeling, and then migrated by Kirchhoff migration to obtain the migration image. In Figure 3b, the resonant multiples, generated by equation 1, are migrated by equation 2. The comparison between Figures 3a and 3b shows that the migration image of resonant multiples has finer resolution. Especially, the top four point scatterers which are a quarter wavelength apart are distinguishable in the resonant multiple migration image.

Figure 3: Migration images of point scatters by migrating (a) primary reflections and (b) resonant multiples. The red dots are the positions of the point scatterers. λ represents the wavelength corresponding to the data wavelet.

THEORY: PREDICTION OF RESONANT MULTIPLES WITH PRIMARY REFLECTIONS

Separation of the resonant multiples is needed before migration. We use ZO primary reflections to predict and so extract resonant multiples. Resonant multiple $d_m(t_m, \mathbf{g}_m)$ recorded at time t_m at $\mathbf{g}_m$ can be predicted by

$$d_m(t_m, \mathbf{g}_m) = \sum d_p(t_p, \mathbf{g}_p), \quad (5)$$

where $t_m = \tau(\mathbf{g}_m; t_p, \mathbf{g}_p)$,

and $d_p(t_p, \mathbf{g}_p)$ is a ZO primary reflection signal recorded at time t_p at $\mathbf{g}_p$. The summation is along an isochron represented by $t_m = \tau(\mathbf{g}_m; t_p, \mathbf{g}_p)$ in the data space. The construction of this isochron in equation 5, that is similar to the migration to zero-offset operation for primaries (Tygel et al., 1998; Schleicher and Bagaini, 2004) can be summarized as a two-step process shown in Figure 4a.

- Step 1: In a ZO primary reflection data set, for each time sample $(t_p, \mathbf{g}_p)$, we build an isochron in the model space. The next step is to use this model space isochron to build the isochron in the data space.
- Step 2: Use the model space isochron as a reflector to forward model resonant multiples. That is, given a

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point at the model space isochron $\mathbf{x}$, calculate the first ray which starts from $\mathbf{x}$ and reflects normally from the free-surface boundary. Based on this ray and the local dip of the isochron at $\mathbf{x}$, find the second ray which obeys the reflection law. The intersection between the second ray and the free-surface boundary is $\mathbf{g}_m$, and twice the total length of the two rays gives the resonant multiple travel time t_m . Repeat this process for all the $\mathbf{x}$ along the model space isochron so that the data space isochron is determined.

In Figure 4b, the red curve is the isochron in the resonant multiple data space, associated with the ZO primary reflection signal represented by the red dot. The result of summing the isochrons associated with the ZO primary reflections is shown in Figure 4c, and the envelop of these of isochrons matches the trajectory of the true resonant multiples.

Figure 5 shows the comparison between the true and predicted resonant multiples based on equation 5. The true multiples are a part of the data generated by finite difference (FD) method with a free-surface boundary condition based on a two-layer velocity model with a dipping interface. The predicted and the resonant multiples match quite well. Notice the wavelets of the predicted resonant multiples are stretched compared to the true multiple wavelets. This stretch is because of the property of the isochron $t_m = \tau(\mathbf{g}_m; t_p, \mathbf{g}_p)$, that is when t_p is perturbed by T_0 , the corresponding perturbation in t_m is longer than T_0 .

NUMERICAL RESULTS

Figure 6a shows a common offset gather (142 m offset) of a marine data set. The events above the blue dashed line are mostly primary reflections, and the events below contain resonant multiple events. A moveout correction and stacking are applied to the primary reflections and to the multiples separately based on their moveout formula. Figure 6b shows the post-stack marine data. Compared with Figure 6a, the signal-to-noise ratio in Figure 6b is strengthened. We then use the post-stack primary reflections in Figure 7b to predict resonant multiples based on equation 5. The prediction results are shown in Figure 7a. This prediction can be used as a reference to identify and to window the resonant multiples in Figure 7b. Also, the predicted resonant multiples say that the third event between the two red dashed lines (indicated by the blue arrow) in Figure 7b is not a resonant multiple.

Figures 8 and 9 show the prediction and migration of resonant multiples in another near-offset (108 m offset) marine data set. The near-offset primaries are used to predict resonant multiples (Figure 8b), and the prediction says that the event indicated by the blue arrow in Figure 8a is a non-resonant multiple event. Migration image of the resonant multiples is shown in Figure 9b. Compared with the post-stack primary image (Figure 9a), resonant multiple image has better resolution.

Figure 4: Illustration of predicting resonant multiples with ZO primary reflections. (a) Transformation from the ZO primary reflection signal recorded at t_p by a trace at $\mathbf{g}_p$ to the resonant multiple signal recorded at t_m by a trace at $\mathbf{g}_m$. (b) The isochron of the resonant multiples (red curve) associated with the ZO primary reflection signal recorded at t_p at $\mathbf{g}_p$ (red dot). (c) A group of isochrones in the resonant multiples (red curves) associated with all of the ZO primary reflections. The envelop of the red curves matches the trajectory of the true resonant multiples.

CONCLUSIONS

We derive the migration formula for surface-related resonant multiples. This migration operator has better resolution than ZO primary reflections migration operator. We also show how ZO primary reflections can be used to predict 1st-order resonant multiples. The prediction of resonant multiples can be used to identify and window the resonant multiples from other events, as shown by a marine seismic data set. The result of another marine data set shows the subwavelength characteristics of imaging resonant multiples.

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Figure 5: Comparison of the predicted and the true resonant multiples in the case of a dipping interface. The black wiggles represent the true resonant multiples, and among them the second event is the 1st-order surface-related resonant multiple (amplitude amplified). The red wiggles represent the predicted resonant multiples.

Figure 7: Comparison between (a) the predicted resonant multiples and (b) the post-stack multiple events. The red dashed lines show that the predicted resonant multiples can be used to window resonant multiples. The blue arrow indicates a non-resonant multiple event.

Figure 6: (a) A pre-stack near offset marine data set. (b) Post-stack data. In panels a and b, events above the blue dashed line are mostly primary reflections, and the events below contain resonant multiples. The blue dashed square sections highlight the better signal-to-noise ratio of post-stack data.

Figure 8: Comparison between (a) post-stack multiple events and (b) predicted resonant multiples from ZO primaries. The prediction in panel b indicates a non-resonant multiple event (blue arrow) in panel a.

Figure 9: Comparison between (a) primary reflections migration image and (b) resonant multiples migration image. Panel b shows better resolution than panel a. The blue arrow indicates migration artifacts caused by non-resonant multiple events.

EDITED REFERENCES

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REFERENCES

- Bleistein, N., J. Cohen, and W. J. Stockwell, 2001, Mathematics of multidimensional seismic imaging, migration, and inversion: Springer Science & Business Media. <http://dx.doi.org/10.1007/978-1-4613-0001-4>.
- Schleicher, J., and C. Bagaini, 2004, Controlling amplitudes in 2.5 d common-shot migration to zero offset: *Geophysics*, **69**, 1299–1310. <http://dx.doi.org/10.1190/1.1801946>.
- Schuster, G. T., and Y. Huang, 2014, Far-field super-resolution by imaging of resonance: *Geophysical Journal International*, **199**, 1–7.
- Tygel, M., J. Schleicher, P. Hubral, and L. T. Santos, 1998, 2.5-d true-amplitude Kirchhoff migration to zero offset in laterally inhomogeneous media: *Geophysics*, **63**, 557–573.